## Analysis of measurements collected in gravity laboratory in Józefosław Observatory during 2007-2010

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## Spring gravimeter LaCoste&Romberg Earth Tide NO. 26 SINCE 2001

BALLISTIC GRAVIMETER FG5 NO. 230 SINCE 2005



## RAW OBSERVATION OF ET GRAVIMETER



## ■ CALIBRATION OF SPRING GRAVIMETER USING AG MEASUREMENTS



## TIDAL PARAMETERS INVESTIGATION



## ATMOSPHERE INFLUENCE ON GRAVITY



## OCEAN LOADING



## HYDROLOGICAL EFFECTS



## BACKGROUND NOISE



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AG measurements

ASTRO-GEODETIC Observatory in Józefoslaw Insur Watazawai is equipped with teo gravimeters for different purposes. Continuously recording LCR ET-25 spring graviteler (since 2022 secent for determination of accurate local tal effects such as almospheric and ocean infuand an analy EQ5 on 210 ballatic analysis ter /einve 2005) is operated periodically - once a Activities

along with FG5 periodically taken measurem

The results for particular series and AG measure-

Figure 2. Scale factor values - upper graph, for sessions of

minimum 2 days length has bigger marks. Number of FGS measurements days and RMS of LCR residuals Tidal parameters determination

MEASUREMENTS with LCR-ET and FG5 pro-

in processing and long series of collected data al-

month. Engineerity meansurance allows us for study non-tidal gravity changes caused mainly by local and continental hydrology. In this paper we present some advantages of using teo types of crash manurements. Outron calibration process the search search have belietin maximular and used for determination of scale factor of script asso for determination of scale tactor of spring anyimeter. On the other hand ballistic anyimete ries of synchronous measurements were used for determination of background noise, atmospheric

(admittance factor) ocean and badminaical al (admittance factor), ocean and rydrological ar-fect on gravity changes. The results from both resimplary is reserved and that areamant is

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Figure 4. Amplitude botors, difference in amplitude factors relative to Waitr Dehart tidel model and phases for main

### Atmosphere influence on gravity

We computed pressure admittance as simple re creasion coefficient on basis of LCR measure



Figure 5. Seasonal variation of atmospheric pressure

### Ocean loading

Tidal gravity parameters in diurnal and semi-In Fig. 4 one can see significant discrepancies be durnal bands are computed using international tween determined and predicted from model body standard data processing techniques. We used tide especially for M<sub>2</sub> constituent. Computing in 40 months (2007-2010) of continuous gravity meadirect effect using most recent ocean models (we surements. The standard deviation of least square adjustment reached 0.98 uzs/x<sup>2</sup>. do not differentiate them here, as they give similar results) greatly reduces this differences.

> lows for investigation in weak environmental sigrais , ressure and ocean loading, betalogical signals. Combining those results with records from different instruments (meteo, GNSS, water table level and soil moisture observations) in Józefoslaw Observatory makes it unique place in Poland for

Figure 6. Phaser plots for residual values (subtracted bod

### Hydrological effects

AG values show periodical variation. Part of sea sonal signal can be explained by loacal and global water storage (Fig. 7).



Figure 7. AG measurements compared to gravity change due continental water storage and local water table level

### Background poise

instrumental noise) on basis of raw observation (1 min sampling). Here we present daily standard de viation from records where tide and polynomial of 9" degree were subtracted



Figure 8. Daily and smoothed (Realify curve) RMS for day

peodetic, peodynamic and peophysics studies.

### Acknowledgments

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